

AMINO ACID N IN HOLOCENE PALEOSOLS AS IDENTIFIED BY IR SPECTROSCOPY AND AMPEROMETRIC TECHNIQUES Teresita Chua-Ona¹, Daniel C. Olk², Najwa Alnsour,³ and Michael L. Thompson¹ Agronomy Department, Iowa State University¹, USDA-ARS National Laboratory for Agriculture and

Environment² and North Carolina State University³

IOWA STATE UNIVERSITY **Department of Agronomy**

INTRODUCTION

Organic matter in paleosols is investigated to infer the mechanisms by which carbon can be preserved over millennial time scales. However, there has been limited research on the nature of nitrogen associated with the stabilized organic matter in paleosols. The most abundant forms of organic N in soils are the amino acids (AA); they are typically bound in soils and may occur as peptides or part of polymeric structures, such as peptidoglycan or glycoproteins. Calderoni and Schnitzer (1984) have reported the abundance of amino acids (AA) in humic materials of Italian paleosols based on their charge properties. Amino acids can also be grouped based on the polarity of the R group attached to their α -carbon atoms. The alkyl group (R) in AA can act as tracer for the assimilation of C that is associated with N.



This study attempts to utilize the characteristics of the R groups attached to AA to track and understand the preservation of N in paleosols. We coupled the pulse amperometric technique with infrared (IR) signatures i identifying how N and consequently C are preserved over millennial scales.

OBJECTIVES

Determine the AA content and distribution based or R groups with age or burial of paleosols

Investigate if IR signals coupled with amino acid identification can provide a better means of understanding the preservation of N as well as C

MATERIALS AND METHODS

I. Study Sites

- Claussen site (along Mill Creek near Lawrence, KS; 740800E, 4326200N)
- Farwell site (along the South Fork of the Big Nemah River in southeast NE; two sub-sites, 64 and 65)
- II. Sample Test Materials
 - Unfractionated soils
 - Clay fractions separated by repeated suspension and exhaustive density separations and then

| State - Carles | | | | 🤝 Feature | Radiocarbon Ag | ge (yr B.P.) Determined on | Wood Charcoal | Fine-grained Deposits | | | | Percent of Tot | tal Amino Acid | N | | | |
|--|--|----------------------------------|---------------------------|-------------------------|------------------------------------|---|---|--------------------------------------|----------------------------------|---|---|--|--|---|---|--|--|
| gure 1 . The cutbank vealing the soil strati | at the Clau graphy and | ussen sit 1 ages. (: | e with its c from Mand | ross-sec lel et. al. | tional vi , 2006) | ew show | n on the | right, | Figure 3. A clay fraction | mino acid c ns. | listribution ba | ased on the attac | ched R groups of C | Claussen unfrac | ctionated s | oils and the | |
| Table 1. Select cheClaussen and Farw | emical and vell sites. | physica | l character | istics of | whole so | oil sample | es at the | | Tab fract | le 2 . Amino tions. | o acid distribu | ution based on th | ne attached R grou | ps in Farwell v | whole soils | and clay | |
| Site | Depth | TC | TOC [§] | TN | pH^{\dagger} | Sand | Silt | Clay | Site | Depth | Total N | | | | Total Amino | AA as fraction | |
| CLAUSSEN | CIII | | g kg | | | | g kg | | 100 | | | | Amino Acid-N | | _ Acid N | of total N | |
| Honey Creek (YHC) | 0-15 | 30 | 14 | 1.2 | 7.5 | 86 | 531 | 383 | | | | Charged R [¶] | Uncharged R [#] | Apolar R [§] | | | |
| Honey Creek (YHC) | 25-40 | 25 | 10 | 1.0 | 7.6 | 91 | 604 | 305 | | cm | | | mg kg ⁻¹ | | | % | |
| Robert Creek (RC) | 82-107 | 13 | 10 | 0.8 | 7.7 | 22 | 575 | 404 | WHO | JLE SOIL S | AMPLES | 206(51) | 145(25) | 127(24) | 570 | 20 | |
| Robert Creek (RC) | 117-132 | 12 | 12 | 1.0 | 7.5 | 11 | 604 | 385 | 03 | 0-20 | 1918 | 290 (31) 199 (55) | 143 (23) | 137(24) | 378 241 | 50 26 | |
| Gunder Paleosol 1 | 104-119 | 5 | 5 | 0.5 | 6.4 | 3 | 651 | 346 | 03 | 20-30 | 1339 | 188 (JJ) 55 (59) | | 09(20) | 05 05 | 20 16 | |
| Gunder Paleosol 1 | 129-144 | 5 | 5 | 0.6 | 6.5 | 3 | 623 | 375 | 05 | 90-155 | 000 | 33 (38) | 28(29) | 12(13) | 93 129 | 10 | |
| Gunder Paleosol 2 | 262-277 | 8 | 4 | 0.4 | 7.6 | 18 | 572 | 410 | 03 | 155-155 | 674 | 04(01) 02(71) | 32 (23) | 22(10) | 138 | 12 | |
| Gunder Paleosol 2 | 287-302 | 10 | 3 | 0.4 | 7.7 | 13 | 562 | 424 | 65 | 175 202 | 720 | os (/1) 61 (64) | 23(19) | 11 (9) 12 (12) | 06 | 20 12 | |
| Gunder Paleosol 3 | 525-530 | 7 | 4 | 0.6 | 7.6 | 13 | 682 | 305 | 03 | 0.20 | 1054 | $\begin{array}{c} 01 & (04) \\ 125 & (54) \end{array}$ | 24 (24) 57 (25) | 12 (12) | 90 220 | 15 | |
| Gunder Paleosol 3 | 550-565 | 6 | 5 | 0.6 | 7.6 | 6 | 671 | 323 | 04 | 0-20 | 1004 707 | 123 (34) | 37(23) | 40 (21) 21 (10) | 250 162 | 22 | |
| Gunder Paleosol 3 | 575-590 | 10 | 4 | 0.5 | 7.5 | 30 | 550 | 419 | 04 | 20-41 | (10) | 89 (33) | 42 (20) | 31(19) | 103 | 10 | |
| | | | | | | | | | 04 | 109-154 | 010 520 | 00 (08) 55 (72) | 21 (22) | 10(10) | 90 76 | 19 | |
| FARWELL | | | | | | | | | 04 CI A | 134-130 | 332 | 55 (72) | 15 (20) | 0 (8) | /0 | 14 | |
| 55 | 0-20 | 24 | 21 | 1.5 | 7.7 | 64 | 680 | 256 | CLA 65 | $\begin{array}{c} 1 \mathbf{F} \mathbf{K} \mathbf{A} \mathbf{C} \mathbf{H} \mathbf{C} \\ 0 2 0 \end{array}$ | 2175 | 301 (60) | 153 (23) | 106 (16) | 650 | 10 | |
| 55 | 20-50 | 17 | 15 | 1.3 | 6.5 | 45 | 625 | 330 | 65 | 0-20 20 50 | 3473 | 316(64) | 133(23) 111(22) | 70(14) | /07 | 15 | |
| 55 | 90-135 | 7 | 6 | 0.8 | 6.1 | 13 | 621 | 366 | 65 | 20-30 | 1004 | 123 (60) | 111 (22) | 70(14) 21(11) | 497 | 10 | |
| 5 | 135-155 | 16 | 13 | 1.0 | 6.3 | 6 | 644 | 350 | 65 | 70-155 125 155 | 1094 | 155 (09) 170 (70) | 40(20) | 21 (11) 27 (11) | 193 | 10 | |
| 5 | 155-175 | 8 | 7 | 0.7 | 6.4 | 3 | 621 | 375 | 65 | 155-155 | 1490 | 1/0 (70) | 49 (19) | 27(11) 22(10) | 234 | 17 | |
| 5 | 175-202 | 9 | 8 | 0.6 | 6.4 | 17 | 644 | 386 | 65 | 133-173 | 1570 | 139(72) | 40(10) | 23(10) | 222 | 10 | |
| | 202-217 | 5 | 5 | 0.5 | 6.6 | 68 | 606 | 326 | 64 | 0.20 | 1413 | 140(70) 164(62) | 40(19) | 22(11) 27(14) | 207 | 15 | |
| 4 | 0-20 | 11 | 9 | 1.1 | 6.9 | 21 | 544 | 435 | 04 64 | 0-20 | 1/10 | 104(03) 128(62) | 57(22) | 37(14) 22(15) | 230 | 15 | |
| 54 | 20-41 | 7 | 6 | 0.7 | 7.3 | 16 | 555 | 429 | 04 | 20-41 | 1090 | 130(02) 141(79) | 30(23) | 33(13) | 120 | 20 12 | |
| j 4 | 109-134 | 6 | 5 | 0.8 | 8.1 | 21 | 627 | 352 | 04 | 109-154 | 1330 | 141 (70) $122 (80)$ | 20(13) | $\begin{array}{c} 11 \\ 0 \\ 6 \end{array}$ | 100 | 12 | |
| 1 | 134-156 | 5 | 4 | 0.7 | 7.8 | 31 | 638 | 331 | 04 Values | in parentheses re | $\frac{1349}{1549}$ | $\frac{123(80)}{144}$ | 21 (14) | 9 (6) | 155 | 11 | |
| /a not available HCl-treated samples 1:1 soil-water | | | | | | | | | ¶ Arg, I # Gly, 7 § Ala, V | Lys, Glu, Asp, His Thr, Ser, Cys, Tyr Val, Pro, Leu, Iso, | Met, Phe | | | | | | |
| aussen Clays, no H | IF treatme | nt ¹⁰⁸ | 2 | Farwell | Clays, r | no HF tre | eatment | 1035 | | Val, Pro, Leu, Iso, | Met, Phe | | | | | | |
| 3700 YHC 0_15 o YHC 25-40 P3 0-15 cm P3 25-40 cm P3 50-65 cm 3000 | cm 163 cm n 2000 avelength (cm ⁻¹ | 1105 2 ₁₄₁₉ 100 | 911 86969 800 0 | 3621 3693 | 3000 | 65 0-20 cm 65 20-50 cm 65 90-135 cr 65 135-155 c 65 155-175 c 65 175-202 c Wavele | 1630 n cm cm 2000 ngth (cm ⁻¹) | 1108 91 8 0 1420 1000 | | Clays at bot acid anions The HF-con cm^{-1} . In add C=C that are increases. A total N in th of the total 1 | h sites reveale (Ellerbrock and centrated who lition, the 160 e conjugated w t Farwell, AA e clay fraction N and was 14- | ed IR signatures nd Gerke, 2004; ole soils demonst 00 cm^{-1} signal ob with the C=O. Well sites, AA cor A ranges from 13 ns. Similarly, the -19% of the total | at 1630 cm ⁻¹ , indic Celi et al., 1998) th rated IR signals in tained for HF-treat ncentrations decrea -30% of total N in e AA in the Clauss N in clays (data n | ative of C=O be nat may be asso about the same ed whole soils whole soils wh en whole soils is ot shown). | onds from ciated with region of can be link period of b ile it is 11- ranged from | carboxylic amides. 1600-1650 ed to the urial 19% of the n 10-40% | |
| Claussen Whole Soils, HF treated | | | | | Farwell Whole Soils, HF treated | | | | | > In both sites, the percentage of the total AA consisting of apolar R groups was higher in modern than buried soils. Conversely, the percentage of the total AA with charged R groups was higher i the buried than the modern soils. | | | | | | | |

freeze-dried

- III. Chemical and Physical Analyses
- Particle size (Gee and Bauder, 1986)
- Total CN analysis via combustion (Vario Micro Cube, Elementar, UK)
- TOC of freeze dried soils following a 24-h 0.5 HCl treatment and 3x water rinses

IV. HF Treatment of Whole Soils

Soils pre-treated with 0.1 M HCl followed by 4x treatment of 5 M HF to dissolve inorganics (Fang et al. 2010) and final water rinses before freezedried

V. Infrared Analysis

- 2 w/w% in IR-grade KBr
- Collected 200 scans from 4000-600 cm⁻¹ via DRIFT using a Nicolet Magna IR spectrometer equipped with a cooled MCT detector
- Reported scans were background corrected and smoothed
- VI. Amino Acid Analysis (Martens and Loeffelmann, 2003)
 - 250 mg soil (<0.1 mm)/200 mg freeze-dried clay with 2 mL 4 M methanesulfonic acid (containing 0.2% tryptamine); soil-acid mixture autoclaved @ 121°C for 16 h (Olk et al. 2008)
 - Digest was diluted and then centrifuged before pH adjusted to 4.0-6.0; brought to final 10-mL volume
 - Ten-fold diluted with water, and filtered through 0.2 µm nylon filter



Charged R - Arg, Lys, Glu, Asp, His • Uncharged R - Gly, Thr, Ser, Cys, Tyr • Apolar R - Ala, Val, Pro, Leu, Iso, Met, Phe



soils from Claussen and Farwell sites.

The qualitative analysis of AA has the potential in elucidating the historical sequence of the preservation of N as well as C in these Holocene paleosols.

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