# An Automated Dual-Head Infiltrometer for Measuring Field Saturated Hydraulic Conductivity **DECAGON DEVICES**

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### Introduction

We developed an automated infiltrometer capable of producing variable hydraulic head conditions without actually varying the water depth and without requiring user intervention. This will enable a user to run multiple infiltrometers at one time getting a better quantification of spatial variability. We used the Reynolds and Elrick (1990) and Nimmo et al. (2009) methods of analysis for steady flow from ponded infiltration, in a single ring, which enables us to accurately calculate the saturated hydraulic conductivity ( $K_{f_s}$ ) of soils in the field. This can replace the method that requires constant monitoring and user intervention to change water levels, and can take hours for the user to complete a single measurement.

# **Design & Performance**

•Peristaltic pump for water supply and volume measurement •Capacitance probe for depth control •Air pressure control to set and change head •Hall effect sensor to measure water pump revolutions •3 heavy steel rings for different flow rates •Quick connect sealing lid •CR3000 Data logger for measurement and control





#### **Dual-head Infiltrometer Equations**

Nimmo et al. (2009) compute the field saturated conductivity from

(1)

 $K_{fs} = i/F$ 

where *i* (cm/s) is the steady (final) infiltration rate (volume divided by area) and F is a function that corrects for sorptivity and geometrical effects. *F* is computed from

## Pressure

Range: -20 to 30 cm H<sub>2</sub>O

**Flow Rates** 20 cm ring Low flow rates 2.5x10<sup>-5</sup> to 9.2x10<sup>-4</sup> cm/sec

15 cm ring Medium low rates 1.48x10<sup>-3</sup> to 4.32x10<sup>-3</sup>cm/sec

10 cm ring High flow rates 2.1x10<sup>-3</sup> to 9.6x10<sup>-3</sup>cm/sec



Figure 1. Internal layout of infiltrometer. On left is original prototype with data logger. On right is second edition with purpose built board.



Figure 2. Testing infiltration on the soccer field.

 $F = 1 + (\lambda + D)/\Delta$ 

#### (2)

where D is the ponding depth (cm), d is the insertion depth of the infiltrometer (cm), b is the infiltrometer radius (cm), and  $\lambda$  (cm) is the reciprocal of the Gardner  $\alpha$  which is a characteristic of the soil and its initial water content. Equations 1 and 2 are originally from Reynolds and Elrick (1990).  $\Delta$  in eq. 2 is 0.993d + 0.578b.

If we know infiltration rate at two ponding depths we can eliminate  $\lambda$  using equations 1 and 2, written for the two depths, to get

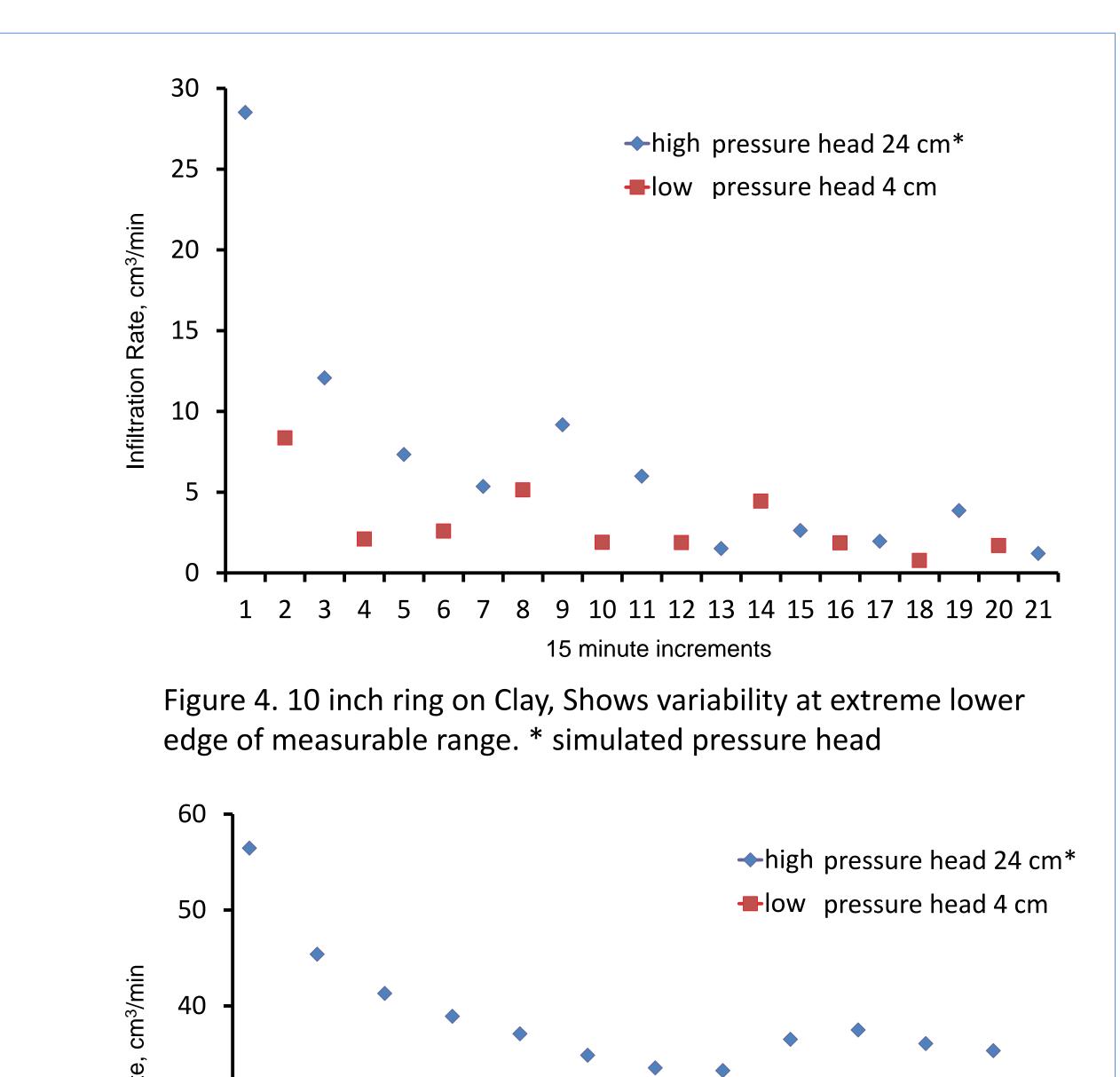
(3)  $K_{fs} = \Delta (i_2 - i_1) / (D_2 - D_1)$ 

 $\Delta$  is a constant for a particular size infiltrometer. For our infiltrometers d = 5 cm and b = 5, 7.5 or 10 cm, so  $\Delta = 7.85$ ,

#### **Field results**

Table 1. Table of infiltration rates for 12 different locations.

Location	Κ <sub>fs</sub>	Ring Size
	cm h <sup>-1</sup>	cm
Clay	0.09	20
Clay	0.61	20
Silt Loam 1	2.20	20
Silt Loam 2	3.32	20
Silt Loam Under Grass	5.33	15
Untilled-Structured Silt Loam	15.55	15
Soccer Field 1	18.47	15
Soccer Field 2	7.78	10
Vantage Sand 1	8.14	10
Vantage Sand 2	9.47	10
Soccer Field 3	21.82	10
Soccer Field 4	34.45	10



9.3 or 10.7 cm. The hydraulic conductivity is therefore just the infiltration rate difference divided by the ponding depth difference multiplied by delta. The infiltration rate is the flow (cm<sup>3</sup>/s) divided by the infiltrometer area. The areas for our infiltrometers is 79, 177 or 314 cm<sup>2</sup>.

References



Nimmo, J. R., K. M. Schmidt, K. S. Perkins and J. D. Stock. 2009. Rapid measurement of field-saturated hydraulic conductivity for areal characterization. Vadose Zone J. 8:142-149. Reynolds, W. D. and D. E. Elrick. 1990. Ponded infiltration from a single ring. I. Analysis of stead flow. Soil Sci. Soc. Am J. 54:1233-1241

