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## Introduction

Soil phosphorus availability is generally assumed to be governed by physico-chemical processes, but biological processes including organic P ( $P_o$ ) mineralization can make important contributions to plant-available P (Bünemann et al., 2012). Mineralized  $P_o$  may be especially important in soils where solution and labile  $P_i$  are low.

Isotopic dilution techniques have been developed to measure  $P_o$  mineralization under steady state conditions (Oehl et al., 2001<sub>a</sub>; 2001<sub>b</sub>). This is done by pairing a short-term Isotopic Exchange Kinetics (IEK) batch experiment, with a longer incubation experiment to separate the physico-chemical and biological processes that govern the amount of isotopically exchangeable  $P_i$  in a soil system.

## Research Objective

To assess gross  $P_o$  mineralization and microbial P uptake in four calcareous forage fields soils along a gradient of Olsen soil test P (STP) concentrations.

## Methods

Soils from two organic and two conventional forage fields in Ontario, Canada, providing a gradient of STP concentrations were selected (Tab. 1) ( $n = 4$  soil samples per field  $\times 4 = 16$ ).

Two main experimental components:

- 1) Isotopic Exchange Kinetics (IEK) experiment: Measured  $^{31}P$  and  $^{33}P$  in solution at 1, 4, 10, 40, 60 and 80 minutes after  $^{33}P$  addition. Parameters of isotopic exchange were determined and used to extrapolate  $E$ -values ( $E_{\text{extrapolated}}$ ).
- 2) 7-day Incubation experiment: Measured  $^{31}P$  and  $^{33}P$  in soil solution and microbial biomass, determined  $E$ -values due to both physico-chemical and biological processes ( $E_{\text{measured}}$ ).

**$E$ -value** – amount of isotopically exchangeable P in a system in a given time period ( $E_t = P_{H_2O} / r_t/R$ ), where  $R$  is amount of radioactivity added initially and  $r_t$  is the amount remaining at time  $t$ .

**Gross  $P_o$  mineralization** =  $E_{\text{measured}} - E_{\text{extrapolated}}$

Tab. 1: Soil properties of the sampled fields.

Soil Description (Field #)	Location (County)	Management Type	Soil Fertility Amendments	Total C (g C kg <sup>-1</sup> )	pH	Olsen STP (mg kg <sup>-1</sup> )	$P_{H_2O}$ (mg kg <sup>-1</sup> )
1	Perth	organic	manure	22.6 <sup>a</sup>	7.4 <sup>a</sup>	2.8 <sup>a</sup>	0.03 <sup>a</sup>
2	Bruce	organic	manure	25.2 <sup>a</sup>	7.2 <sup>a</sup>	5.5 <sup>ab</sup>	0.07 <sup>a</sup>
3	Perth	conventional	manure and fertilizer	23.6 <sup>a</sup>	7.3 <sup>a</sup>	6.8 <sup>b</sup>	0.10 <sup>a</sup>
4	Bruce	conventional	manure and fertilizer	25.9 <sup>a</sup>	7.5 <sup>a</sup>	11.2 <sup>c</sup>	0.24 <sup>b</sup>



Fig. 1: Labelling and weighing out soil for the incubation experiment.

## Results

Gross  $P_o$  mineralization accounted for a mean of 35% of isotopically exchangeable P across all four field soils (Tab. 2)

After 8 days, > 20% of the  $^{33}P$  label was found in the microbial biomass P pool in 11/16 soils.

Microbial  $^{33}P$  uptake was most rapid for soils with the lowest available  $P_i$  (Fig. 2).

As microbial P concentrations were constant, the data suggests accelerated P cycling (turnover) at the lowest solution P concentrations < 0.1 mg P kg<sup>-1</sup>.

Gross  $P_o$  mineralisation was positively related to water soluble  $P_i$  concentrations (Fig. 3); however, this relationship deserves further study, as 3/16 soil samples with low solution  $P_i$  concentrations may have been underestimated due to very rapid microbial P uptake.

Tab. 2. Mean extrapolated ( $E_{\text{ext}}$ ) and measured ( $E_{\text{meas}}$ ) isotopically exchangeable P ( $E$ ) values determined for the four field soils. Gross organic P ( $P_o$ ) mineralization (min) is calculated by difference.

Soil	$E_{\text{ext}}$ (mg kg <sup>-1</sup> )	$E_{\text{meas}}$ (mg kg <sup>-1</sup> )	Gross $P_o$ min. (mg kg <sup>-1</sup> )
1	12.8 <sup>a</sup>	23.7 <sup>a</sup>	10.9 <sup>a</sup>
2	17.5 <sup>a</sup>	24.7 <sup>a</sup>	7.2 <sup>a</sup>
3	19.8 <sup>ab</sup>	30.4 <sup>a</sup>	10.2 <sup>a</sup>
4	28.9 <sup>b</sup>	42.6 <sup>b</sup>	14.1 <sup>a</sup>
<b>Average</b>	<b>19.8</b>	<b>30.3</b>	<b>10.5</b>

## References

Bünemann et al. (2012). Soil Biol Biochem. 51: 84-95.  
Oehl et al. (2001<sub>a</sub>). Soil Sci Soc Am J. 65: 780-787.  
Oehl et al. (2001<sub>b</sub>). Biol Fert Soils 34:31-41.

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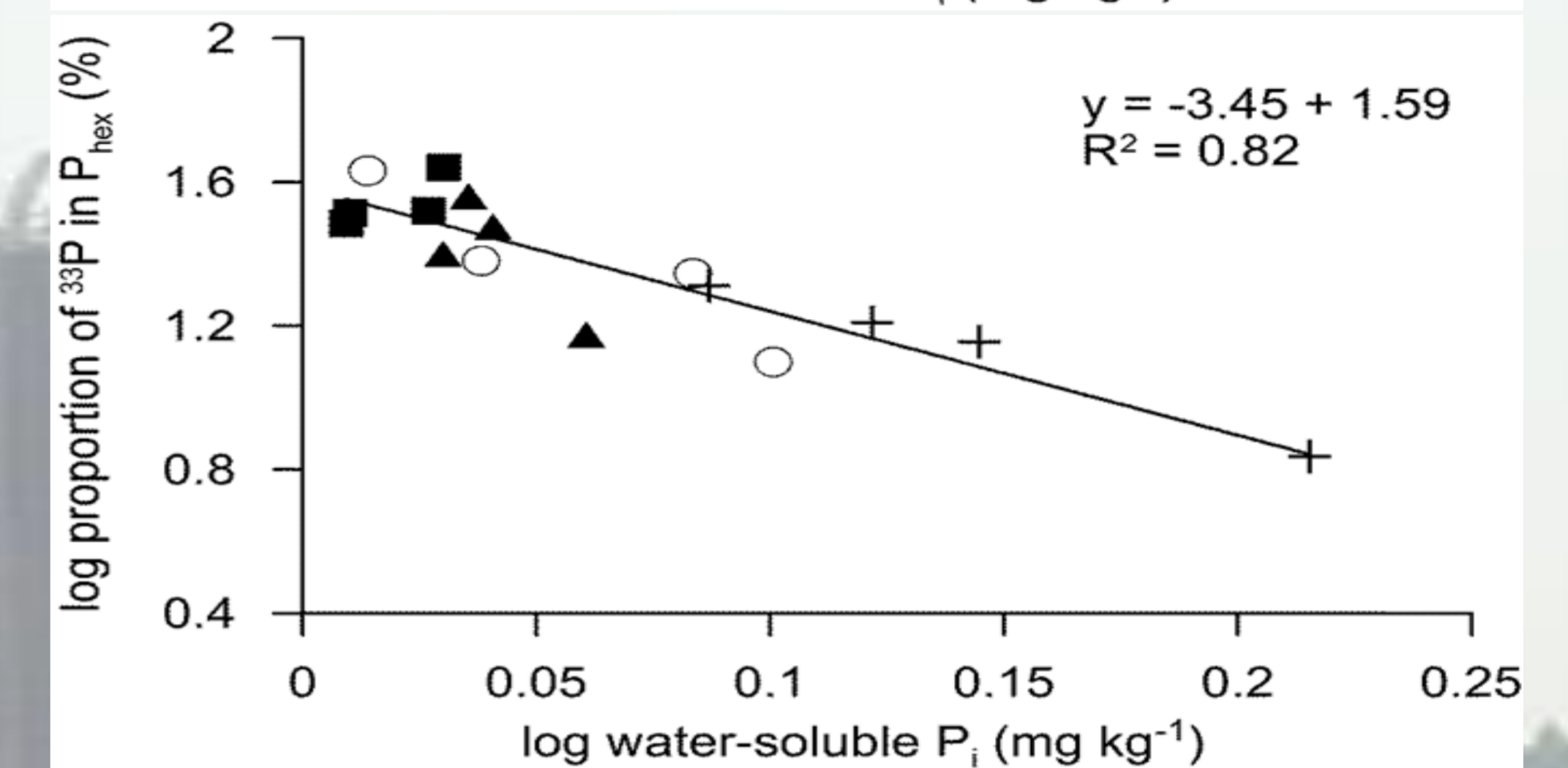
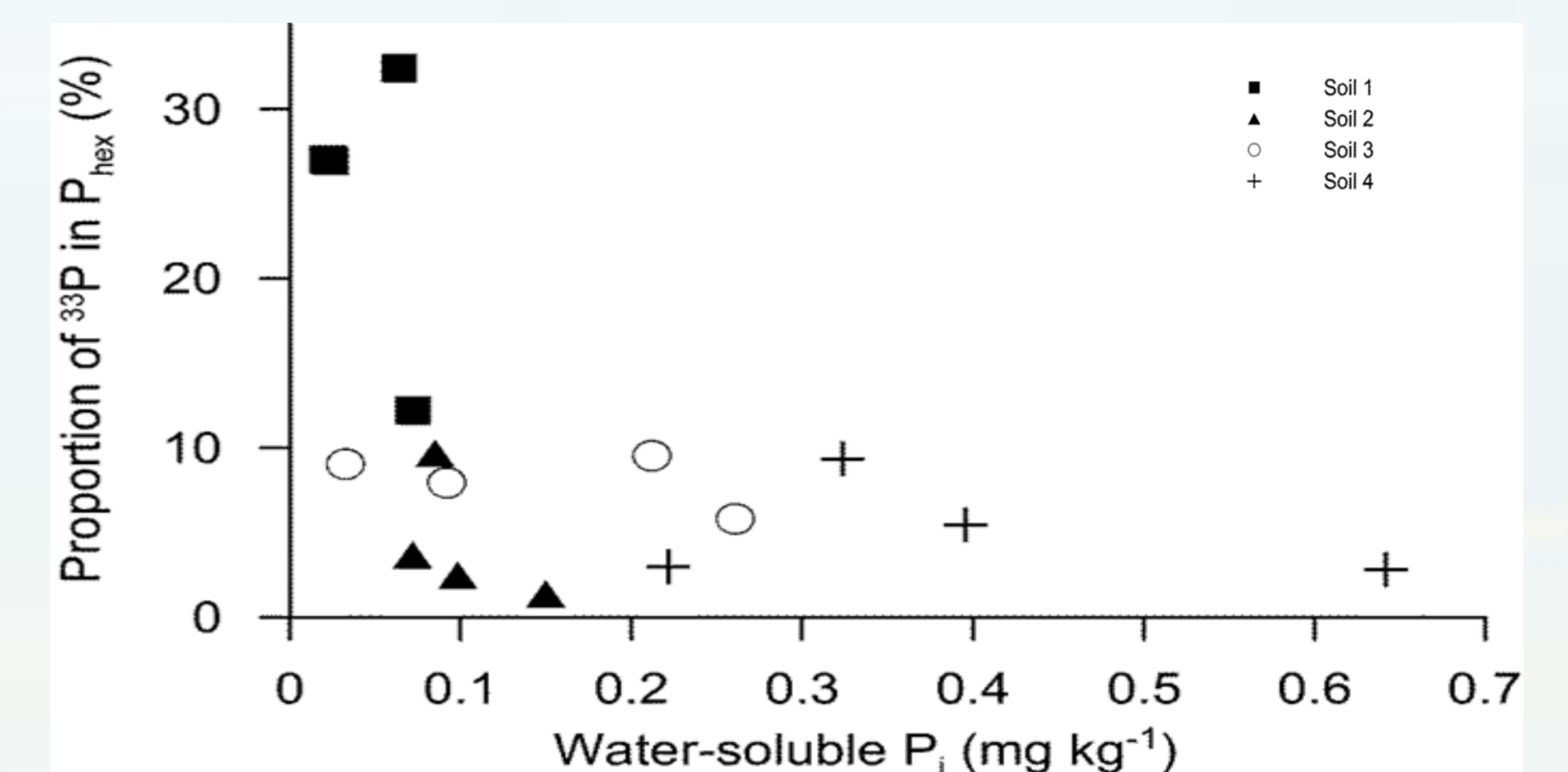


Fig. 2. Proportion of  $^{33}P$  label added found in the microbial biomass P pool ( $P_{\text{hex}}$ ) one (top) and eight (bottom) days after soil labelling.

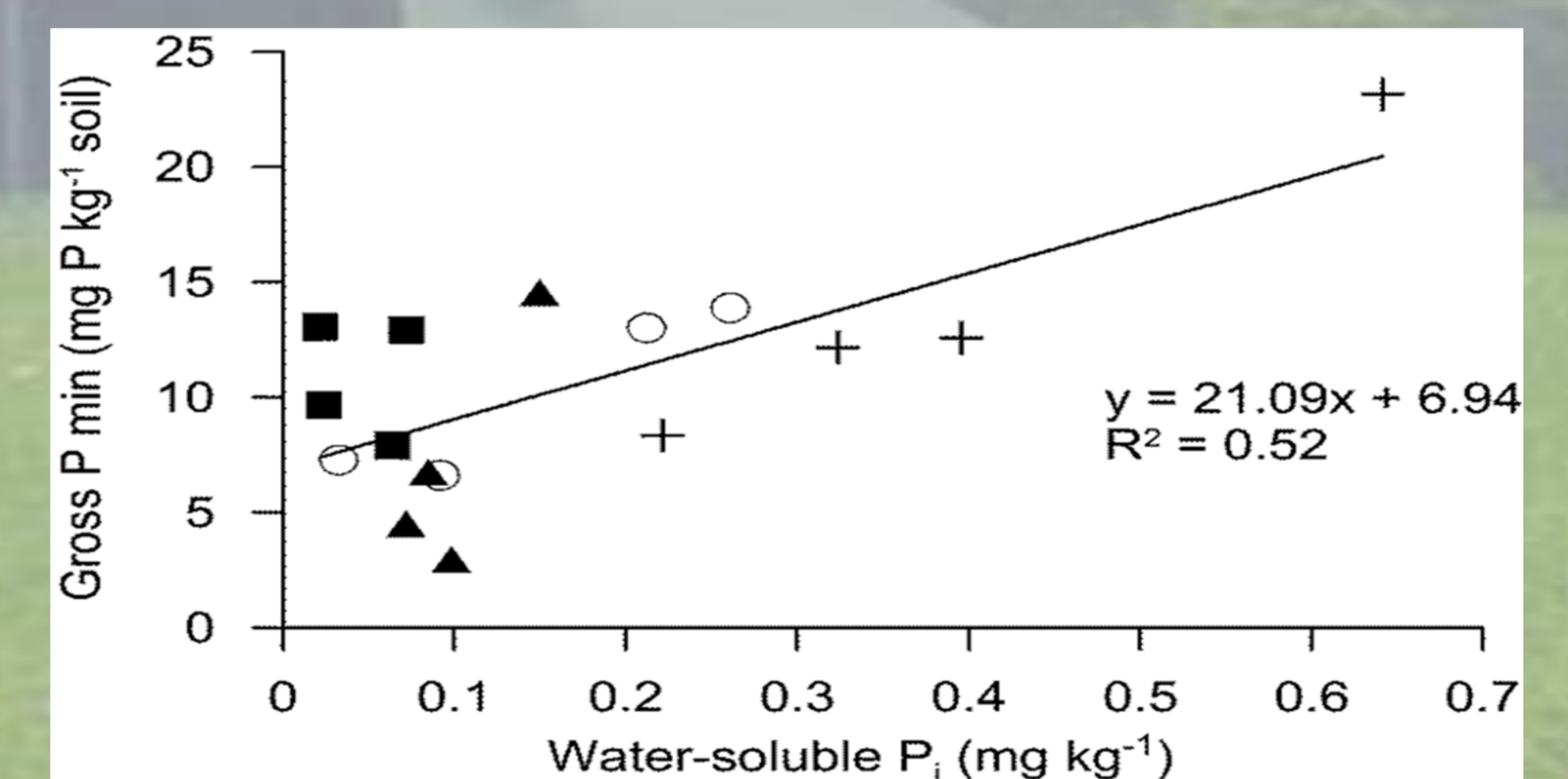


Fig. 3: Gross organic P mineralization (min) after eight days of soil labelling vs. soil water soluble P concentration.

## Conclusions

Gross  $P_o$  mineralization accounts for a significant amount of isotopically-exchangeable P in forage soils.

Rapid uptake of  $^{33}P$  into microbial biomass P pool indicates important P cycling occurs under steady-state conditions. The data suggest microbial P turnover is accelerated under low conditions of low solution  $P_i$  concentrations.

Methodological challenges including the use of high P-fixing soils with low water soluble  $P_i$  concentrations need to be further addressed when using these methods.

## References

Bünemann et al. (2012). Soil Biol Biochem. 51: 84-95.  
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