

Evaluation of Water Vapor Sorption Hysteresis in Soils: The Role of Organic Matter and Clay



UID: 92016

Emmanuel Arthur¹(Emmanuel.Arthur@agro.au.dk), Markus Tuller², Per Moldrup³, Lis W. de Jonge¹

¹Dept. of Agroecology, Aarhus University, Denmark. ²Dept. of Soil, Water and Environ. Sci. The Univ. of Arizona, USA.

³Dept. of Civil Engineering, Aalborg University, Denmark.



Introduction

- Water sorption hysteresis (H) is the difference exhibited in the relationship between the water content (w) of a soil and the corresponding water potential/relative humidity (RH) obtained by wetting or drying
- Extensive literature exist on causes and quantification of H for soil water potential range from 0 to -1.5 MPa but information on H is limited for water potentials < -10 MPa
- Consideration of H in the range from -10 to -480 MPa is crucial for modeling physical and biological soil processes

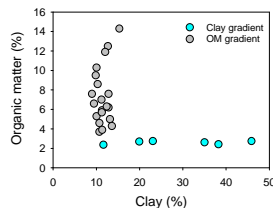
Objectives

- Assess and compare recently developed methods for quantifying water vapor sorption hysteresis in soils and pure clays for the water potential range of -10 to -480 MPa
- Investigate the role of organic matter (OM) and clay content and type on water vapor hysteresis

Methods

- Investigated Samples
Five pure clays: Kaolinite, Illite, Vermiculite, Halloysite, Montmorillonite

- Two groups of soils
- (i) Six soils with clay gradient (11-46%) and OM-2.6%,
 - (ii) 20 soils, OM gradient (3-15%) and clay content-11%.



- Sorption Isotherm Measurements
Wetting and drying isotherms measured with Vapor Sorption Analyzer (-10 to -480 MPa; pF 5.0 to 6.6; RH 3 to 93%)



- Hysteresis Quantification Methods
(i) Based on number of molecular layers (n) from a modified BET (MBET) isotherm equation^{1,2}

$$w = \frac{RH(1 - RH)^n}{[(k_1 + k_2RH)(1 - RH)]}$$

Model parameters: k_1, k_2, n
 n = molecular layers in multilayer

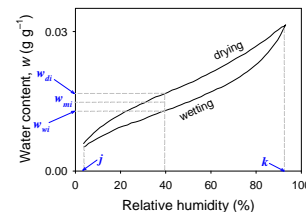
Model fitted separately to wetting and drying curves to obtain " n " and calculate H_1

$$H_1 = \frac{n_{wetting} - n_{drying}}{n_{wetting}}$$

- (ii) Average Degree of Hysteresis³, D_h

$$H_2 = \frac{\sum_{i=j}^{i=k} \frac{w_{di} - w_{wi}}{w_{mi}}}{k - j + 1}$$

10 data points between 3% and 93%RH selected for calculating H_2

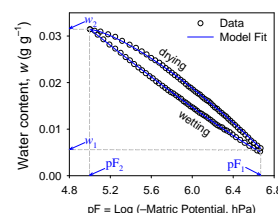


- (iii) Single parameter non-singularity model⁴, SPN

$$pF = pF_1 - (pF_1 - pF_2) \left(\frac{w_1 - w}{w_1 - w_2} \right)^N$$

Model fitted separately for wetting and drying curves to obtain ' N ' and calculate H_3

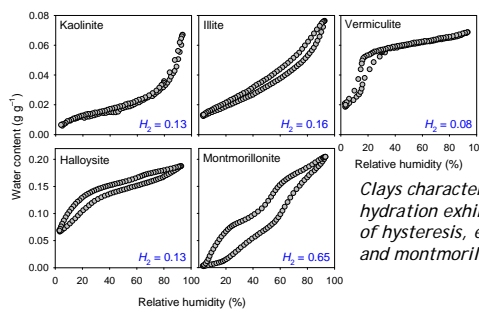
$$H_3 = \frac{N_{wetting}}{N_{drying}}$$



Results

Pure clays

MBET- n and SPN methods were unable to capture hysteresis
 D_h method accurately described hysteresis (H_3)



Clays characterised by interlayer hydration exhibit larger degree of hysteresis, e.g. cf. kaolinite and montmorillonite

Soils

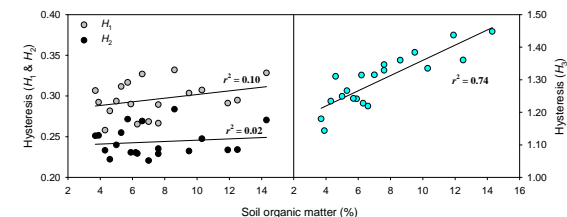
All 3 methods successfully quantified H for both groups of soils
Clay gradient soils

Clay%	12	20	23	35	38	46
H_1	0.22	0.25	0.24	0.22	0.22	0.28
H_2	0.21	0.24	0.23	0.21	0.21	0.25
H_3	1.22	1.19	1.16	1.33	1.37	1.46

H_1 and H_2 : no clear relationship with clay content
 H_3 : increases with increasing clay content

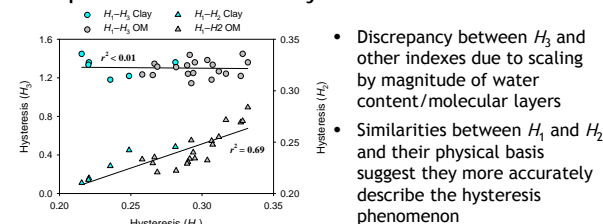
Organic matter gradient soils

Relationship between soil OM and the three hysteresis indexes (H_1, H_2, H_3)



- No clear effect of OM on H_1 or H_2
- For H_3 , large contents of organic matter associated with greater degree of hysteresis

Comparison of the three hysteresis indexes



Trend of larger H_3 values with increasing clay or OM could be a reflection of increasing water content, not actual hysteresis

Conclusions

- All three methods accurately captured hysteresis for soils; but for pure clays, only the D_h method was appropriate
- For pure clays, extent of interlayer hydration determines the degree of hysteresis
- For soils, OM and clay contents showed no clear effect on H

Acknowledgments

The study was financed by the Danish Council for Independent Research | Technology and Production Sciences via the project Water Vapor Sorption Isotherms as Proxy for Soil Surface Properties (DFP -4184-00171)

References

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