Application of Shortwave Infrared Imaging for Estimation of Soil Hydraulic Properties Ebrahim Babaeian¹, Morteza Sadeghi², Mohammad R. Gohardoust¹, Emmanuel Arthur³, Scott B. Jones², and Markus Tuller¹

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Introduction

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- Soil hydraulic properties (i.e., soil water characteristic and hydraulic conductivity) are key for modeling water flow and transport of solutes, heat and gases in saturated and unsaturated soils.
- Current laboratory and field techniques for hydraulic property measurements are expensive, time-consuming, and impractical for large scale applications.
- We present a rapid laboratory proximal sensing method for estimation of soil hydraulic properties (SHPs).
- 2. HYRUS 2D Numerical Modeling and Parameter Optimization
- Water imbibition was described based on Richards' equation considering van Genuchten-Mualem hydraulic functions (van Genuchten, 1980):

$$\frac{\partial \theta(h)}{\partial t} = \frac{1}{r'} \frac{\partial}{\partial r'} \left[r' K(h) \frac{\partial h}{\partial r'} \right] + \frac{\partial}{\partial z} \left[K(h) \frac{\partial h}{\partial z} + K(h) \right] - S(h)$$

Results

SWIR imaging-based surface soil water dynamics

- SWIR imaging provides high resolution maps of soil water content during water imbibition.
- With the advancing



The new method infers SHPs from SWIR reflectance imaging of soil moisture during water imbibition into dry soil in conjunction with a physical reflectance model (Sadeghi et al., 2015) and inverse numerical modeling of Richards' equation with HYDRUS 2D (Šimůnek et al., 2008).

Methods

1. Experiments

- Water imbibition experiments with a surface emitter with constant discharge rate $(0.003 \text{ cm}^3 \text{ s}^{-1})$ were conducted.
- Soil surface reflectance evolution was imaged with a benchtop SWIR line scan camera.





With: $S_e = \frac{\theta - \theta_r}{\theta_s - \theta_r} = \left[1 + |\alpha h|^n\right]^{-m}$ and $K(h) = K_s S_e^l \left[1 - \left(1 - S_e^{1/m}\right)^m\right]^{-m}$

HYDRUS 2D was utilized for water flow modeling (*Šimůnek et* al., 2008). The flow domain was defined as axisymmetrical with three observation nodes located at r' = 1.25, 2.5, and 3.75 cm:



All SWIR image-derived soil water contents for pixels with Euclidian distances of 1.25, 2.5 and 3.75 cm from the emitter were averaged and considered as the observed soil water

wetting front, three zones including saturated zone, unsaturated (transition) zone, and dry zone were observed.

Estimation of $\theta(h)$ and K(h)

• The estimated (converted) drying $\theta(h)$ and wetting K(h) (i.e., K_s) showed reasonable agreement with lab-measured $\theta(h)$ and K_s . However, there exist some discrepancies at the dry-end ($\theta < 0.2$), particularly for finer textured soils, which can be potentially attributed to inaccuracy of the used retention model at the dry-end (*Tuller et al.*, 1999).





Soil Water Content Simulations

- A reasonable agreement is observed between the imaging-based estimated and HYDRUSsimulated water contents with respect to distance from the emitter, particular for finetextured soils where RMSE ranges between 0.025 and 0.067 for sand, and between 0.008 and $0.018 \text{ cm}^3 \text{ cm}^{-3}$ for clay.
- A reasonable agreement is also observed between imaging-derived and HYDRUS-based simulated water contents with respect to time, especially for the first two nodes, indicating reasonable performance of the optimization method to minimize the objective function.





• Reflectance maps were transformed to surface water content maps with a nonlinear model introduced in Sadeghi et al. (2015):



- θ : volumetric water content (cm³ cm⁻³) θ_{s} : saturated volumetric water content (cm³ cm⁻³) **R**: average reflectance (-) *r*: transformed reflectance (-) r_d and r_s : transformed reflectance at dry and saturation (-) σ : shape parameter (-)
- Soil reflectance decreases with the advancing wetted region and increasing water content providing an opportunity to monitor surface water content dynamics and thereby retrieving soil hydraulic properties.
- The nonlinear function was fitted to θ/θ_s values determined from water mass balance and corresponding r values to optimize r_d and r_s parameters.





contents at the 3 nodes and used in the objective function.



HYDRUS 2D was coupled with a global optimization method (Simulated Annealing) in MATLAB to optimize hydraulic parameters α , n, θ_r and K_s through minimizing the objective function (ϕ):

$$\varphi(b) = \sum_{j=1}^{N} \sum_{i=1}^{M} \left[O(r'_j, t_i) - S(r'_j, t_i, b) \right]$$

- *b* : vector of the optimized parameters *N* and *M* : the number of estimations and observations in a particular estimation set $O(r'_{i}, t_{i})$: observed water content at time *i* for the *j*th measurement set at location *r* $S(r'_{i}, t_{i}, b)$: corresponding simulated water content for the vector b
- Search space and initial values for the van Genuchten (1980) hydraulic parameters in the global optimization algorithm.

Parameter	Search space		Initial values (NRCS Soil Survey database)				
	Min.	Max.	Sand	Sandy loam	Loam	Silt loam	Clay
α (cm ⁻¹)	0.005	0.150	0.145	0.075	0.036	0.020	0.008
<i>n</i> (-)	1.001	4.00	2.68	1.89	1.56	1.41	1.09
θ_r (cm ³ cm ⁻³)	0.01	0.10	0.045	0.065	0.078	0.067	0.068
K_s (cm s ⁻¹)	1e-10	0.012	0.00825	0.00123	0.00029	0.00012	0.00005

- The simple hysteresis model of Kool and Parker (1987) was applied to evaluate estimated wetting parameters against independently measured soil water characteristic data.
- The flowchart illustrates the SHP estimation framework.



Conclusions

- The method allows the determination of the soil water characteristic and hydraulic conductivity function within only a couple of days, which is an advantage when compared to time consuming standard laboratory methods.
- Obtained results for soils spanning a wide textural range are in reasonable agreement with independently measured (i.e., Tempe cells and WP4-T Dewpoint Potentiameter) soil water characteristics.

References

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